Glacial Climate in the Tropics:

Ca Isotopes as a Recorder for Past Sea Surface Temperatures

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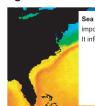
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Introduction

We present palaeo-sea surface temperature (SST) changes from the Equatorial Atlantic based on $\delta^{44}\text{Ca}$ of the planktonic foraminifera *Globigerinoides saculifer*. The frequently discussed question of the amount of tropical temperature change in the upper water column at the Holocene-Last Glacial Maximum (LGM) boundary gave a strong impetus for the current work.



Significance



Sea surface temperature is among the most important variable for Earth's climate system. It influences

sea-air gas exchange global heat transfer evaporation-precipitaion patterns primary biological production continental ice masses deep sea circulation

Calcium isotope basics

Thermal Ionisation Mass Spectrometry (TIMS)

⁴³Ca-⁴⁸Ca double spike technique

 $\delta\text{-notation:}\quad \delta^{44}\text{Ca} = \quad \left[\frac{^{44}\text{Ca}/^{40}\text{Ca}_{\text{sample}}}{^{44}\text{Ca}/^{40}\text{Ca}_{\text{normal}}} - 1 \right] * 1000$

Standardisation: $\Delta(SRM 915a-Seawater) = -1.88 \pm 0.04$

 Δ (SRM 915a-CaF₂) = -1.44 ± 0.04

Magnitude of Cooling in the Tropics

Termination I (Holocene-LGM)

SST _{CLIMAP}	1-2°C	global	faunal assemblages	CLIMAP Members 1981
SST	3°C	tropical oceans	δ18O, alkenone	see W.Broecker 1996
			,	
SST _{GCM}	3.3°C	tropical Atlantic	GCM	Ganopolski 1998
SST	4°-5°C	SW of Barbados	Sr/Ca (corals)	Guilderson et al. 1994
Temperature	4°-6°C	global, Brazil	snowline, noble gas	Stute et al. 1995, Rind et al. 1985

Eastern Equatorial Atlantic (GeoB1112)

Equatorial Pacific

SST _{δ44Ca}	2.5-4.0°C	our study
SST _{Mg/Ca}	3°C	Nürnberg et al. 2000

 SST_{Mg/Ca}
 3°C
 Lea et al. 2000

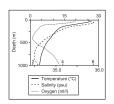
 SST_{Mg/Ca}
 3.5-4°C
 Visser et al. 2003

Conclusion

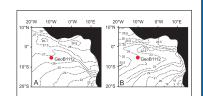
- δ44Ca in *G. sacculifer* is a promising sea surface temperature-proxy
- A possible key to estimate more climate relevant variables
- ullet Overall glacial-interglacial amplitude of $\delta^{44}\text{Ca}$ is 0.6-0.8 permil
- Intermediate temperature changes at Terminations I and II (2.5 to 4°C)
- SST-records (δ^{44} Ca, Mg/Ca) lead the δ^{18} O signal:

Warming predates melting

Oceanographic background of core GeoB1112

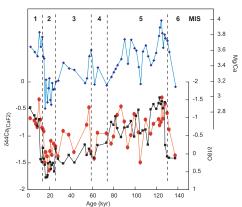


Core GeoB1112 is located in the low productive South Equatorial Current and was recovered from the Guinea Basin from 3125m water depth (5°46.7'S, 10°45.0'W). Average sedimentation rate is 3.0 cm/kyr.

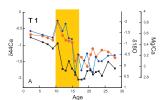


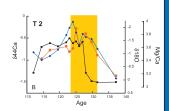
Surface water temperatures in the equatorial Atlantic for February (A) and July (B). Note that the eastern equatorial region is relatively warm in February, that is, temperatures follow seasons of the southern hemisphere.

Down-core variations on fossil G. sacculifer (GeoB1112): A multi-proxy comparison in the tropical East Atlantic



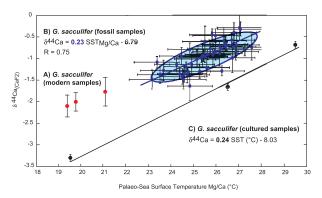
The δ^{44} Ca (red) and corresponding Mg/Ca ratios (blue, Nürnberg et al. 2000) of *G. sacculifer* from sediment core GeoB1112 versus time. Both records show glacial/interglacial fluctuations predicting higher SSTs during interglacial and lower SSTs during glacial periods. The oxygen isotope record also derived from *G. sacculifer* is added for a multi-proxy comparison. It is obvious that the δ^{44} Ca and Mg/Ca records are leading the δ^{18} O signal. This means that warming occurs prior to ice melting by several thousend years. Dashed lines mark the transition between marine isotope stages.





G. sacculifer Ca isotope (orange), Mg/Ca (blue) and oxygen isotope (black) records for GeoB1112 plotted versus time for MIS 1 and 2 (A) and MIS 5e and 6 (B). For both terminations, in equatorial Atlantic the sea surface temperature warming occurs several thousand years before the decrease in ice volume reflected by the oxygen isotope record. Pronounced temperature increase to interglacials (f.e. MIS 5e) seems to be characterised by a double peak.

$\delta^{44}\text{Ca-Thermometry}$ based on G. sacculifer



A) Preliminary δ44Ca ratios of modern G. sacculifer B) The SST(Mg/Ca)-cross-calibrated thermometer based on fossil G. sacculifer support the refined calibration (A) in slope and absolute temperatures. C) The δ44Ca ratios of G. sacculifer cultured under controlled temperature conditions (Nägler et al. 2000). The temperature gradients are identical. However, there is an apparent offset of approximately 3.5°C in absolute temperature. (A) and (B) tend to predict lower sea surface temperatures than (C). The most probable reason is that cultured samples are biased.

Reference

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