

Assessing carbon dynamics in Amazonia with the Dynamic Global Vegetation Model LPJmL – discharge evaluation

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Introduction

The Amazon Basin covers an area of approximately 6 million km² and plays a vital role in global water and carbon cycles (SCHLESINGER & MELACK 1981, BROWN et al. 1993, MEYBECK 1993, ANDREAE et al. 2002). The Amazon River discharges nearly 15% of the earth's fresh water (GAILLARDET et al. 1997).

In this huge system the transport and the retention of dissolved and particulate carbon are influenced by climatic variations (e.g. El Niño) and climate change (AALTO et al. 2003, LEWIS et al. 2004, SCHOENGART et al. 2004) as well as local and regional land use management (FEARNSIDE & BARBOSA 1998, BERNARDES et al. 2004, THOMAS et al. 2004). The main components of the carbon flow through Amazonia are net ecosystem productivity in terrestrial ecosystems (affected by climate, CO₂, and land use), transformation of organic matter in the terrestrial and the aquatic ecosystem, and net export of carbon from land to river water body (affected by litter production, erosion, and flooding regime) (FEARNSIDE 1997, HEDGES et al. 2000, WATERLOO et al. 2006).

A substantial, globally significant carbon flow exists between the atmosphere, land ecosystems, and aquatic ecosystems in the Amazon basin, probably leading to a major, yet poorly constrained carbon export to the Tropical Atlantic. One estimate of the annual export of carbon to the ocean is 40 Tg C/yr (MOREIRA-TURCQ et al. 2003). The total basin evasion has been estimated to be as much as 470 Tg C/yr (RICHEY et al. 2002).

To understand the complex interactions between land, river, climate change, and land use change, we use the dynamic global vegetation model LPJmL, a terrestrial biosphere model (SITCH et al. 2003, BONDEAU et al. 2007). For our purpose, we modify the model and link the terrestrial processes with a hydrological runoff model to include the riverine fluxes. This enables us to assess the influence of changes in land use and climate on regional and global carbon fluxes.

Key words: Amazon, dynamic global vegetation model, riverine carbon fluxes;

Study site

The simulated Amazon basin is enclosed by coordinates 4.75°N to 20.25°S and 50.25°W to 79.25°W. This geographical dimension is defined by a hydrological runoff model (JACHNER et al. in press) and corresponds with published data (BIRKETT et al. 2002, COE et al. 2002). Amazonia is characterized by annual flooding, mainly driven by precipitation (SCHOENGART et al. 2004). During the high-water stage from March to June, the river rises up to 15 m (JUNK & WEBER 1996), connecting the riverine and terrestrial system.

Methods

We use LPJmL, a process-based dynamic global vegetation and hydrology model. It simulates growth, abundance, and population dynamics of natural vegetation and agricultural crops at a resolution of 0.5° (latitude/longitude).

The simulations were driven by monthly climate data (mean temperature, precipitation, number of wet days, sunshine hours) based on CRU05 (NEW et al. 2000, ÖSTERLE et al. 2003). Soil texture data as well as atmospheric CO₂ concentration were used, as in SITCH et al. (2003).

The model was brought into equilibrium by running 1000 years of spin-up, followed by the transient simulation from 1901 to 2003. The hydrology model was adjusted for the Amazon with a velocity of 0.25 m/s, a value in the dimension of published data (BIRKETT et al. 2002, HU et al. 2004). We compared the simulated and the observed discharge at 3 gauging stations in the Amazon basin using data from the 'River Discharge Database' (SAGE). The only stations in the Amazon Basin are Alta Mira at the Rio Xingu (3.25°S; 52.25°W; years 1976–1979), Porto Velho at the Rio Madeira (8.75°S; 63.75°W; years 1969–1979), and Obidos at the mainstem of the Rio Amazon (1.75°S; 1.75°W; years 1971–1983). Our comparison used the simulated data from years for which we had observed data for each site.

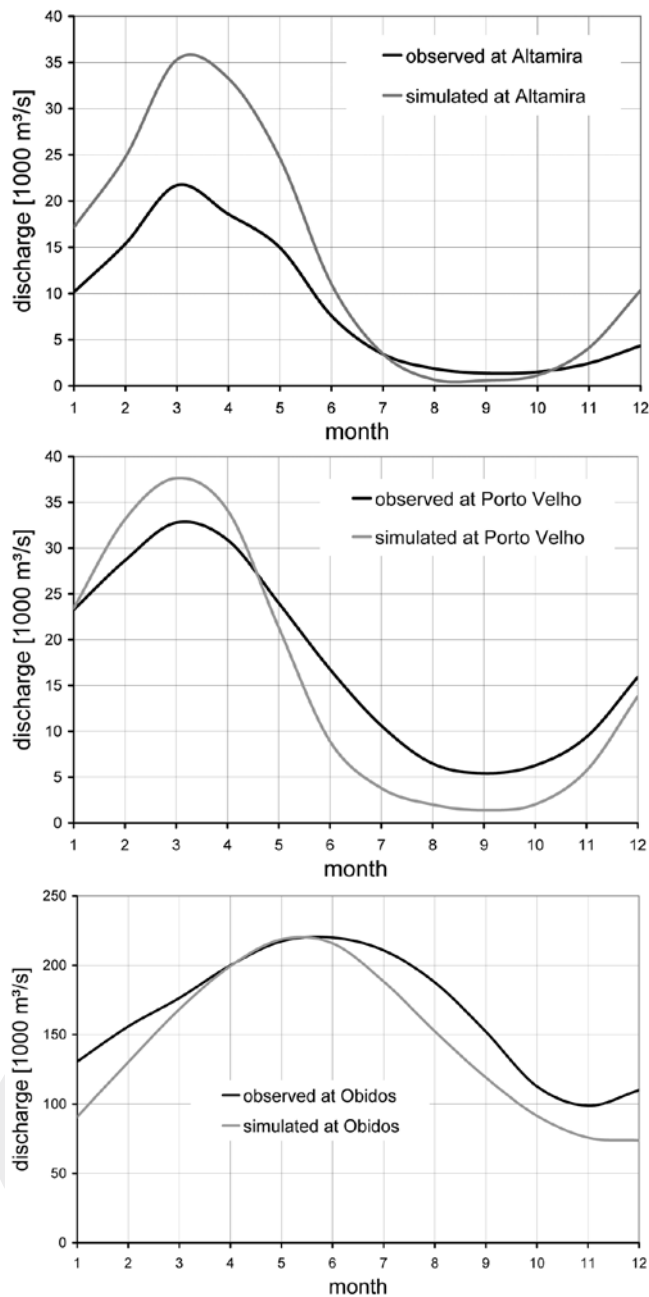


Fig. 1. Comparison of observed and simulated monthly discharge at Altamira (a), Porto Velho (b), and Obidos (c).

Results

We compared the annual and monthly discharge at the 3 stations (Fig. 1).

Alta Mira: The annual total simulated discharge is 1.63 ± 0.11 times the observed discharge. Both simulated and observed high water peaks are in March. The simulated high water discharge is 1.63 times the observed discharge.

Porto Velho: The annual total simulated discharge is 0.90 ± 0.12 times the observed discharge. Both high water peaks are in March and the simulated discharge in March is 1.15 times the observed discharge.

Obidos: The annual total simulated discharge is 0.88 ± 0.06 times the observed. The observed high water peak is in June, but the simulated peak is in May. The simulated discharge at the high water peak is 0.99 times the observed discharge.

For all stations the annual total simulated discharge is 1.00 ± 0.28 times the observed, and the simulated high water discharge is 1.26 ± 0.33 times the observed.

Discussion

Our purpose is to understand the carbon fluxes in the Amazon basin, including fluxes between the terrestrial and the riverine part. Therefore, it is of special importance to reproduce the seasonal dynamics and the height of flooding.

The differences between the annual discharge for all station is very low. Our intention was to optimize the annual discharge, and therefore we had to accept the differences in discharge during the high water peak (factor ranging from 0.99 to 1.63). The differences in the time of high water peak for Obidos between the simulated (May) and the observed (June) data can be explained by the nature of the data. Both datasets are monthly data. A weekly resolution (for which no data are available) could indicate a peak between May and June and would thereby support our findings.

The high water peak in Altamira and Porto Velho is simulated earlier than the peak in Obidos. This matches the observed pattern where the tributary discharge is out of phase with the mainstem (RICHEY et al. 1986).

Our results support that LPJmL can reproduce the pattern and can therefore be applied to riverine carbon flux modelling. Prospective results of LPJmL can include prognosis of the “tipping point” in Amazonia (SCHELLNHUBER et al. 2005).

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